Kiruna-type apatite-iron ore in Svalbard: the evidence from main and accessory minerals

Maraszewska, M.^{1,*}, Manecki, M.¹, Majka, J.^{1,2}, Ziemniak, G.¹, Czerny, J.¹, Schneider, D.³, Faehnrich, K.¹, Ofierska, W.¹, Myhre, P.I.⁴, Barnes, Ch.¹, Włodek, A.¹

¹Department of Mineralogy, Petrography and Geochemistry, AGH University of Science and Technology, Kraków, Poland ² Department of Earth Sciences, Uppsala University, Villavävagen 16 SE-75266 Uppsala Sweden ³Department of Earth and Environmental Sciences, University of Ottawa, Canada ⁴Norwegian Polar Institute, Tromsø, Norway * E-Mail: marramaraszewska@gmail.com

Apatite-iron oxide deposits (so called "Kiruna-type" iron ores) are phosphate-rich magnetite-hematite ores associated with magmatic or volcanic rocks. They are originally known from Paleoproterozoic rocks of the Baltic Shield in Sweden, and subsequently recognized in formations of various age in many other places around the world. Recent studies suggest a complex, combined magmatic and post-magmatic origin for iron oxides, and phosphates that crystallized from evolved siliceous melt, highly enriched in fluids and REE. This is reflected in the trace element chemistry of the ore as well as accessory minerals, for example enrichment in V, Mn, Cr and depletion in Ti and Ni in iron oxides, and high concentration of F and REE in phosphates. However, hydrothermal post-magmatic overprint obscures the primary geochemistry.

Small apatite-iron oxide ore lensoidal bodies were found in the northeast Prins Karls Forland, an island of the Svalbard Archipelago. They occur locally within a narrow (~1 km wide) shear zone between two distinct units: (1) the amphibole facies Pinkie Unit comprising metapelites, calc-silicate rocks and marbles to the east and (2) the greenschist facies metasediments of the Grampianfjella Formation to the west. Hydrothermally altered and slightly metamorphosed cumulate gabbro bodies, several meters in size, occur in the vicinity of the ores (Maraszewska et al. 2016). The gabbros contain relatively high amounts of ilmenite with secondary titanite rims and xenomorphic chlorapatite. Both the gabbros and the ores are hosted by tectonized metasediments of the ductile-brittle shear zone.

Iron oxides in the ores are present as magnetite and hematite. The latter is probably the product of magnetite alternation resulting from the increase of oxygen fugacity during the hydrothermal event. They are accompanied by chlorapatite and cut by quartz-chlorite veins. The structures and textures of the ore vary from augen structure to massive and even skeletal. In the augen type, strongly deformed porphyroblasts of magnetite are surrounded by hematite-dominated matrix. In the massive ore, magnetite form porphyroblasts whereas specularite hematite is developed along extension cracks and veins. In skeletal form, significantly altered type, dendritic iron oxide aggregates are overgrown by secondary titanite and iron hydroxides.

Electron microprobe chemical analysis shows that most of iron-oxides are depleted in Ti (\sim 0.01-0.1 wt%) as well as Zn, Mn, and Ni. The content of Al and Cr is moderate, whereas V (up to 0.35%) and Si (0.1-0.4%) occur in relatively higher amounts. Projection of the results on Al+Mn vs. Ti+V as well as Ti vs. Ni/Cr plots indicates that these magnetites could be related to the Kiruna-type deposits (Fig. 1).



Fig. 1. A – Plot of Ti versus Ni/Cr in magnetite. B – Al+Mn versus Ti+V in magnetite discrimination diagram. Large pink dots represent magnetite samples from Prins Karls Forland on Svalbard. All other data after Dare et al. (2014) in A and after Dupuis et al. (2011) in B

Accessory REE phosphates, such as monazite and xenotime form ahedral inclusions, several μ m in size, in magnetite crystals, less commonly in silicate matrix. Monazite is Ceand Nd-enriched (25.5-29 wt% and 10.5-13.5 wt%, respectively). Y-xenotime is generally homogenous. However, some zones show slight enrichment in LREE. Uranium and thorium content in both minerals does not exceed 0.5 wt%. Such chemistry of these phases is typical in Kiruna-type deposits (Jonsson et al. 2016).

Co-existence of iron oxides with apatite and REE-phosphates as well as geochemical signature of magnetite indicate that the Prins Karls Forland ores belong to the Kirunatype. However, such deposits contain usually fluorapatite and are associated with intermediate or felsic magmatic rocks (e.g. Harlov et al. 2002; Jonsson et al. 2016). The Prins Karls Forland ores seem to be associated with more mafic rocks. The presence of Ti-Fe-oxides and chlorapatite in these gabbros strengthens the possible genetic link. It is the first occurrence of such ores in the High Arctic.

Acknowledgments: This work is partially funded by AGH research grant no 11.11.140.319.

References:

- Dare SAS, Barnes SJ, Beaudoin G, Meric J, Boutroy E, Potvin-Doucet C (2014) Trace elements in magnetite as petrogenetic indicators. Miner Deposita 49:785–796
- Dupuis C, Beaudoin G (2011) Discriminant diagrams for iron oxide trace element: finger-printing of mineral deposit types. Miner Deposita 46:319–335
- Harlov DE, Andersson UB, Forser HJ, Nyström JO, Dulski P, Broman C (2002) Apatite-monazite relations in the Kirunavaara magnetite-apatite ore, northern Sweden. Chem Geol 191:47–72
- Jonsson E, Harlov DE, Majka J, Högdahl K, Persson-Nilsson K (2016) Fluorapatite-monazite-allanite relations in the Grängesberg apatite-iron oxide ore district, Bergslagen, Sweden. Am Mineral 101:1769–1782
- Maraszewska M, Manecki M, Czerny J, Schneider D, Myhre PI, Faehnrich K, Barnes Ch (2016) Metagabbro associated with the shear zone on Prins Karls Forland (Svalbard, Arctic). Geophysical Research Abstracts, EGU General Assembly 17-22 April 2016, Vienna